Multiscale Modeling of CO2 Flow and Storage in Pre-Salt Reservoirs: Perspectives and Challenges

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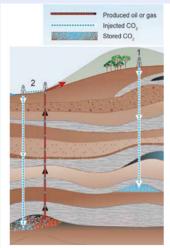
> and Marcio Murad LNCC/MCTIC





OUR MAIN TARGET

Feasibility of Vuggy Carbonates (Pre-Salt) as Potential Storage Sites for CO2 Storage



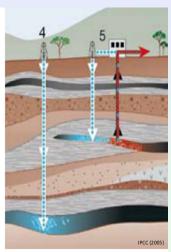
Oil fields

- 1- Depleted reservoirs (gas/oil)
- 2- Enhanced oil recovery

- 1km _2km

Saline Aquifers;

3- Deep unused saline water-saturated reservoir rocks.



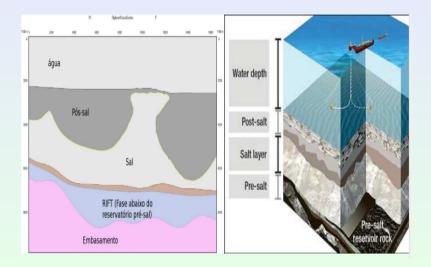
- Coal layers.
- 4- Deep Unmineable coal
- 5- ECBM Recovery

ABC

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Possible Choices for Storage Sites

Carbonate Underneath the Salt Layer or Rock Salt

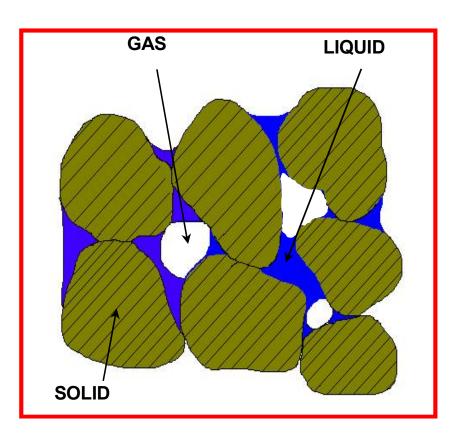


OUTLINE OF THE LECTURE

Pros and Cons

- Storage Site: Carbonate underneath the Salt Layer
 - Review of CO2-properties at the Subsurface
 - CO₂ movement and Driving Forces
 - Thermodynamics
 - Geochemistry
- Review of the Main Trapping Mechanims in the Subsurface
 - Stratigraphic, Residual,
 - Solubility, Mineral
- Drawbacks: (Ongoing Research)
 - Geomechanical Issues
 - Apperance of High Permeability Pathways Conduits
- Storage Site: Salt Layer

• Concluding Remarks on Feasibility of CO2 Storage in the Pre-Salt



The **species** are:

- mineral (-) : main mineral
- water (w) : as liquid or evaporated in the gas phase
- air (a) : dry air, as gas or dissolved in the liquid phase
- chemical species : interacting (reactive) species

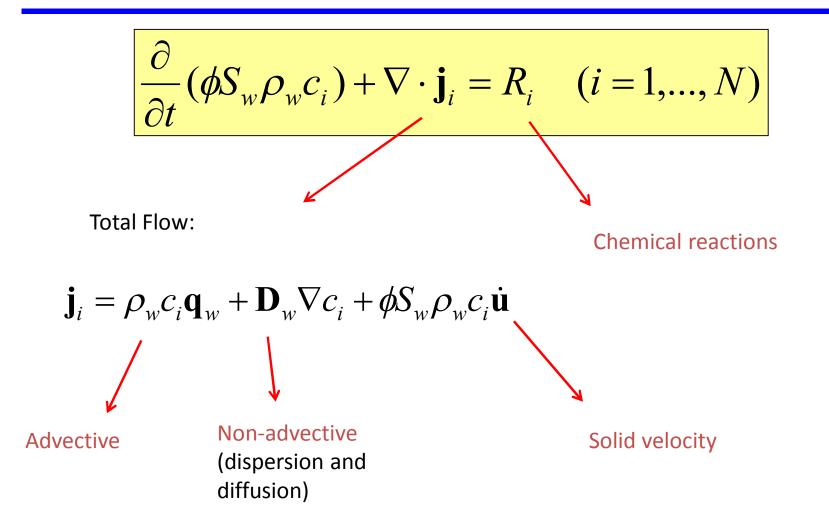
The three phases are:

- **gas** (g) : mixture of dry air and water vapour
- liquid (I) : water + air dissolved + dissolved chemical species
- solid (s) : main mineral + absorbed cations + precipitated minerals





Reactive transport equations





Reactive transport equations

$$\frac{\partial}{\partial t}(\phi S_w \rho_w c_i) + \nabla \cdot \mathbf{j}_i = \mathbf{R}_i \quad (i = 1, ..., N)$$

□ CHEMICAL INTERACTION OF *N* INTERACTING SPECIES

- Slow reactions: kinetics controlled
- Fast reactions: equilibrium controlled
- PHENOMENA CONSIDERED
 - Homogeneous reactions
 - Aqueous complex formation
 - Acid/base reactions
 - Oxidation/reduction reactions
 - Heterogeneous reactions
 - Cation exchange
 - Dissolution/precipitation of minerals (equilibrium and kinetics)
 - Other reactions
 - Radioactive decay
 - Linear sorption

Reactive transport equations



$$\frac{\partial}{\partial t}(\phi S_l \rho_l c_i) + \nabla \cdot \mathbf{j}_i = R_i \quad (i = 1, ..., N)$$

□ CHEMICAL INTERACTION OF **N** INTERACTING SPECIES

Slow reactions: kinetics controlled

 Rate of species production in kineticscontrolled reactions

$$\mathcal{V}_{m} = A_{m}k_{m}\left|\Omega_{p}^{r}-1\right|^{n}$$
$$\Omega_{p} = \frac{Q_{m}}{K_{m}} ; \qquad Q_{m} = \prod_{j=1}^{N_{c}}a_{j}^{v_{mj}}$$
$$k_{m} = k_{25}\exp\left[\frac{-E_{a}}{R}\left(\frac{1}{T}-\frac{1}{298.15}\right)\right]$$

□ Fast reactions: equilibrium controlled

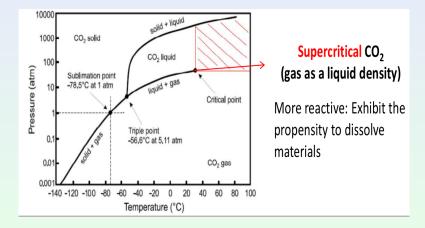
 A chemical equilibrium model is uses based on the minimization of Gibbs free energy

$$\begin{aligned} \underset{n_{j}^{c}, n_{i}^{x}}{minimize} \quad G &= \sum_{j=1}^{N_{c}} \mu_{j}^{c} n_{j}^{c} + \sum_{i=1}^{N_{x}} \mu_{i}^{x} n_{i}^{x} \\ n_{j}^{U} &= n_{j}^{c} + \sum_{i=1}^{N_{x}} \nu_{ij} n_{i}^{x} \quad (j = 1, ..., N_{c}) \\ n_{i}^{x} &\geq 0 \quad (i = 1, ..., N_{x}) \\ n_{j}^{c} &\geq 0 \quad (j = 1, ..., N_{c}) \end{aligned}$$

- Newton-Raphson algorithm
- Lagrange multipliers to incorporate the restrictions of the system

Thermodynamics: Supercritical State

Temperature: 31.1*C*, Pressure 7.38 MPa: Geothermal gradients 25C/KM, z = 800m.



• In the long-term behaves as a separate phase with much lower Volume

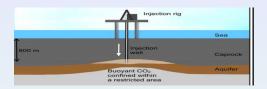
ABC

Desirable Properties for a Successful Injectivity

- Efficient Trapping Mechanics
- Host formation not impermeable
- Deep enough to maintain Co2 in superctitical state
- High areal extent of the cap rock
- Avoidance of Geomechanical Structural Damage
- After Long Term Dissolution/Precipitation remains in a Mineralized state



Trapping Mechanics i) Structural/stratigraphic Free Gas Phase



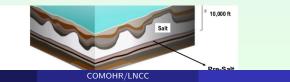
Two-Phase Flow in Porous Media. Buoyant Forces

$$\mathbf{v}_{T} = -KK_{r}(S)(\nabla P - \rho \mathbf{g})$$

$$\phi \frac{\partial S}{\partial t} + \nabla \cdot (\mathbf{v}_{T}f(S)) = 0$$

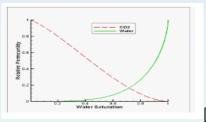
$$f_{=}f_{w}(1 - \alpha(Kg(\rho_{w} - \rho_{g})))$$

Salt (Halite and Dolomite) – Excellent non-reactive Impermeable Medium.



ABC

Also referred to as capillary trapping



[Relative Permeability Curves

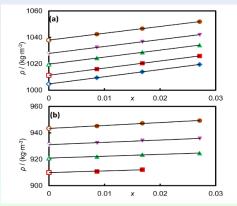
$$oldsymbol{v}_{\mathcal{T}} = -rac{K}{\mu} K_r(S) (
abla P -
ho oldsymbol{g}), \qquad \phi rac{\partial S}{\partial t} +
abla \cdot (oldsymbol{v}_{\mathcal{T}} f(s)) = 0$$

• A fraction of the CO₂ is left behind as a disconnected phase

Trapping Mechanics iii) Solubility in the Aqueous Phase

Increase in Brine Density

 $\rho = \rho_{water}(1 - X) + X \rho_{CO2}, \qquad X - -\text{mass fraction}C_{O2}$



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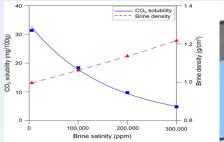
Top T= 296 K:

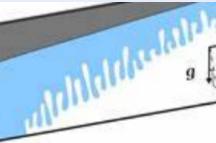
Bottom T= 449 K

ABC

Trapping Mechanics iii) Solubility in the Aqueous Phase

Increase in Brine Density – Gravitational stability, $\rho_w = \rho_w(C_{CO2})$





- Moves in the Opposite DIrection of the Free Gas.
- Sinks towards the bottom of the host formation
- Mass fraction Convective-Diffusion Reaction Equation

$$\phi \frac{\partial \phi(\rho_w X_i)}{\partial t} + \nabla \cdot (\rho_w X_i \mathbf{v}_D) = \nabla \cdot (D_i \nabla (\rho_w X_i)) + F_i, \qquad i = CO_2, \text{ water}$$

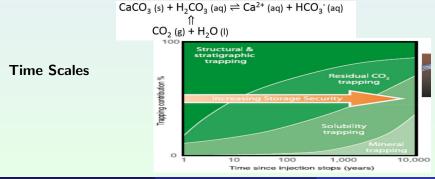
Trapping Mechanisms: iv) Mineral Trapping

• Solubility $\rightarrow \rightarrow$ Acidification $\rightarrow \rightarrow$ Weak carbonic acid

 $CO_2 + H_2O = HCO_3^- + H^+$

(water acidification)

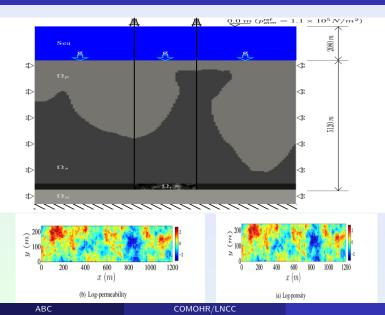
Trigger Dissolution – Precipitation Geochemical Reactions



ABC

Drawbacks: Structural Geomechanics

i) Integrity of the Cap Rock



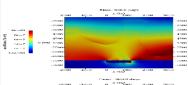
Hydro-Mechanical Multi-Physics Coupled Model

- Two-Phase Flow, Elasticity of the Host Rock,
- Viscoelasticity of the Cap Rock Salt

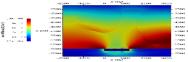
Fixed Stress Split: Creep of the Rock Salt) div $\boldsymbol{\sigma}_{e} = -\rho_{s}^{DOMO}\mathbf{g}$ div $\boldsymbol{\sigma}_{e} = -\rho_{e}^{POS}\mathbf{g}$ $egin{aligned} & m{\sigma}_e = \mathcal{C} \mathbb{E}_e + m{\sigma}^0 \ & \mathbb{E} =
abla^s \mathbf{u} \end{aligned}$ $\sigma_{e} = \mathcal{C}\mathbb{E}_{e} + \sigma^{0} \quad \text{in } \Omega_{POS}$ $\mathbb{E}_{e} = \frac{1}{2} \left(\nabla \mathbf{u} + (\nabla \mathbf{u})^{T} \right)$ $\mathbb{E}_{e} = \mathbb{E} - \mathbb{E}_{v}$ $\mathbf{S} = 2G\mathbf{P}_{dev}\mathbb{E}_{e}$ in Ω_{SALT}
$$\begin{split} \dot{\mathbb{E}}_{\mathbf{v}} &= \dot{\gamma} \frac{\partial \sigma_{V}}{\partial \sigma} = \dot{\gamma} \sqrt{\frac{3}{2}} \frac{\mathbf{S}}{\|\mathbf{S}\|} \\ \dot{\gamma} &= \mathcal{E}_{R}^{*} \left(\frac{\sigma_{V}}{\sigma_{R}}\right)^{N} \quad \sigma_{V} = \sqrt{\frac{3}{2}} \|\mathbf{S}\| \end{split}$$
div $\boldsymbol{\sigma}_{e} = \nabla \boldsymbol{\rho} - \left((S_{w} \rho_{w} + (1 - S_{w}) \rho_{o}) \phi + \rho_{s}^{RES} (1 - \phi) \right) \mathbf{g}$ $oldsymbol{\sigma}_e = \mathcal{C}\mathbb{E}_e + oldsymbol{\sigma}^0, \qquad \mathbb{E}_e = rac{1}{2} \Big(
abla \mathbf{u} + (
abla \mathbf{u})^T \Big)$ $\beta \frac{\partial p}{\partial t} + \operatorname{div} \mathbf{v}_{Dt} = -\beta \frac{\partial \tilde{\sigma}}{\partial t}$ in Ω_{RES} $\mathbf{v}_{Dt} = -\lambda_t(S_w)K(\phi)(\nabla p + \mathbb{E})$ $\frac{\partial(S_w\phi)}{\partial t} + \operatorname{div}\left(\lambda_w(S_w)(\mathbf{v}_{Dt} - \boldsymbol{\eta}k_{ro})\right) = 0$

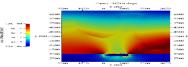
Saturation and Von-Mises Stress

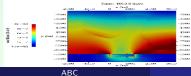
Elastic x Viscoelastic



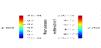
1.1







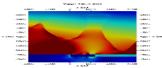


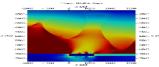


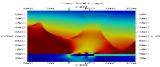


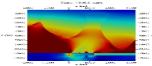


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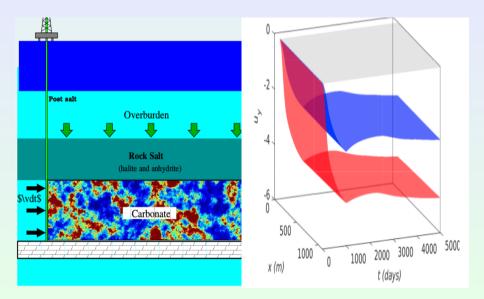








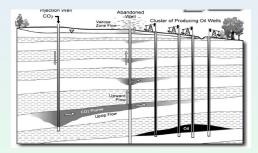
Surface Uplift



Structural Geomechanical Issues

ii) Leakage through Abandoned Wells

- Saline Aquifers; Do not suffer from this possibility
- Oil fields. Anthropogenic activity.



- Preexisting Abandoned wells penetrating caprock
- Potential leakage high-permeability pathways for buoyant CO₂

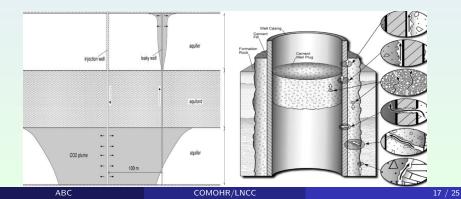
Structural Geomechanical Issues

ii) Leakage through Abandoned Wells

Two-Phase Flow (Celia et al 2013)

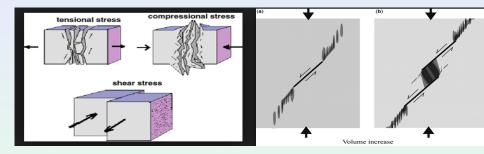
$$\phi \frac{\partial \rho_g S_g}{\partial t} - \nabla \cdot (\rho_g \lambda_g \mathbf{K} (\nabla P - \rho_g \mathbf{g})) = 0$$

$$\phi \frac{\partial S_w}{\partial t} - \nabla \cdot (\lambda_w \mathbf{K} (\nabla P - \rho_w \mathbf{g})) = 0$$



Structural Geomechanics

iii) Fault Activation due to Fluid Over-Pressurization



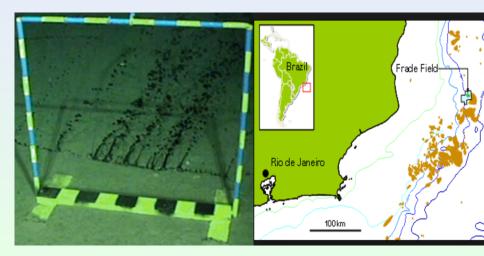
Localized Leaking: Plasticity. Mohr Coulomb Function. $\dot{\mathcal{E}}_{p} = \lambda \partial F(\sigma) / \partial \sigma$



Structural Geomechanics

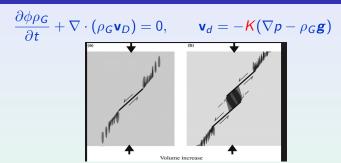
iii) Fault Activation

• Frade Field Oil Spill Incident

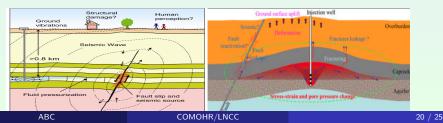


iii) Fault Activation

Dilatancy: Increase in Permeability



Injection-Fault-Activation \rightarrow Ground Surface Uplift and Induced Siesmicity



Structural Geomechanical

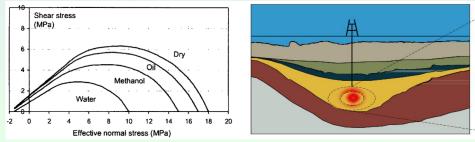
iv) Water Weakening. Loss of Injectivity

Injectivity Index

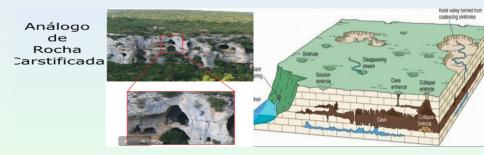
$$I = \frac{\text{Injection Flow Rate}}{P_{inj} - P_{res}}$$

Low *pH* triggers dissolution/precipitation reactions new the injector Geomechanics: Poroplasticity: Decrease strength of the rock bonds.

$$\nabla \cdot \boldsymbol{\sigma} - \nabla P = 0, \qquad \boldsymbol{\sigma} = C(\boldsymbol{\mathcal{E}}(u) - \boldsymbol{\mathcal{E}}_p)$$
$$\dot{\boldsymbol{\mathcal{E}}}_p = \lambda \, \partial F(\boldsymbol{\sigma}, S) / \partial \boldsymbol{\sigma}$$

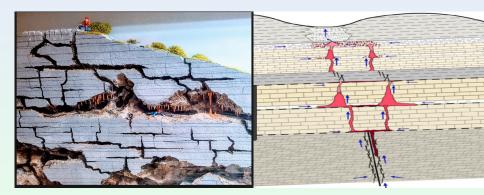


• Fractured Corridors. Enlarged due to Dissolution Collapse

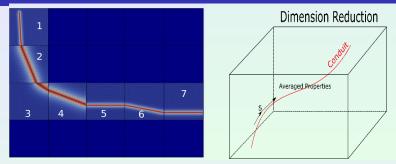


Presence of a Cave Network

Karst Conduits



Coupled 3D - 1D Model



$$\beta_r \frac{\partial P_r}{\partial t} + \nabla \cdot \mathbf{v}_d = \mathbf{K} \mathbf{I} (P_r - P_c) \delta_\Delta \quad \text{in} \quad \Omega$$
$$\mathbf{v}_d = -\mathbf{K} \nabla P_r$$

Coupled System

$$\beta_c \frac{\partial P_c}{\partial t} + \frac{dv_c}{ds} = -\frac{KI}{A_c}(P_r - P_c)$$
$$v_c = -K\frac{dp_c}{ds} \quad \text{in} \quad \Gamma_c$$

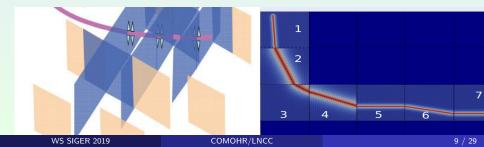
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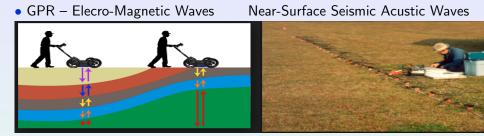
Summary - High-Fidelity (Fine-Scale) Model

$$\beta_{m} \frac{\partial P_{m}}{\partial t} - \nabla \cdot (K_{m} \nabla P_{m}) = \mathbf{K} \mathbf{I} (P_{m} - P_{c}) \delta_{\Delta} \quad \Omega \subset \mathbb{R}^{3}$$
$$\beta_{f} \frac{\partial P_{f}}{\partial t} - \nabla_{\tau} \cdot (K_{f} \nabla_{\tau} P_{f}) = (\mathbf{v}_{m}^{+} - \mathbf{v}_{m}^{-}) \cdot \mathbf{n} \quad \Gamma \subset \mathbb{R}^{2}$$
$$\beta_{c} \frac{\partial P_{c}}{\partial t} + \frac{dv_{c}}{ds} = -\frac{\mathbf{K} \mathbf{I}}{A_{c}} (P_{r} - P_{c})$$
$$v_{c} = -K_{c} \frac{dp_{c}}{ds} \quad \text{in} \quad \gamma_{c} \quad \mathbb{R}^{1}$$

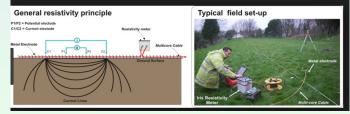
 δ_{Δ} – Dirac line source



III Incorporate Outcrop Data

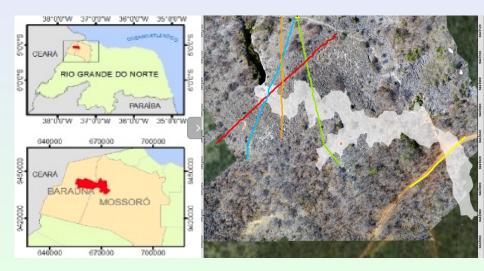


• Electro-Resistivity - Electric Current - Electrodes



• - Joint Work with F Hilario Bezerra and F Pinheiro (UFRN)

Outcrop FURNA FEIA



RADAR FACES

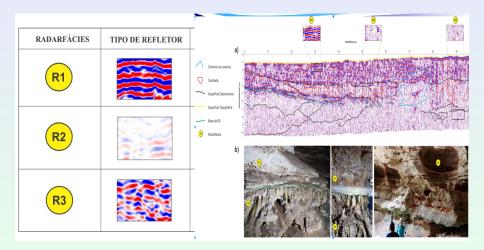
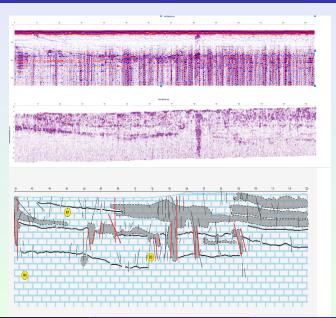


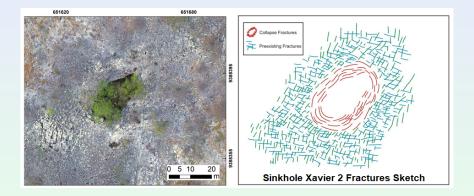
IMAGE PROCESSING



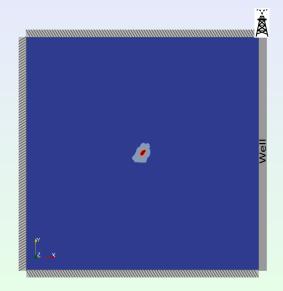
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Sinkhole Xavier 2



Sinkhole Xavier 2



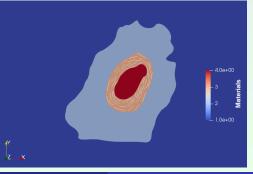
Sinkhole Xavier 2 - Materials

Material	Description	Permeability [m ²]
1	Intact Rock	$K_x = K_y = 2.96076 \times 10^{-13}$
2	Intact Rock + Micro fractures	$K_{\rm x} = 3.07392 imes 10^{-13}$
	(estimated via ODA)	$K_y = 3.08085 imes 10^{-13}$
3	Intact Rock + Micro fractures	$K_{\rm x} = 3.07392 imes 10^{-13}$
	(estimated via ODA) + discrete fractures	$K_y = 3.08085 imes 10^{-13}$
4	Sinkhole	$\mathcal{K}_x = \mathcal{K}_y = 1.0 imes 10^{-8}$
White Lines	Discrete Fractures	$K_f = 8.333 imes 10^{-10}$
		(aperture $= 1.0 imes 10^{-4}$ m)

Estimating ODA:

Micro-Fractures Permeability $K_f = 8.333 \times 10^{-10} \text{ m}^2$

Micro-Fractures Aperture $d = 1.0 \times 10^{-4} \text{ m}$



Parameters	Values
Domain	500m imes 500m
Rock Permeability (Horizontal)	$K_x = 3.2433 \times 10^{-13} \text{ m}^2$
Rock Permeability (Vertical)	$K_y = 3.2815 \times 10^{-13} \text{ m}^2$
Initial Pressure	56 MPa
Well Pressure	55 MPa
Sinkhole Pressure	55 MPa
Karst Index	$7.370811 \times 10^{-13} \text{ m}^3$
Total Simulation Time	$4 imes 10^5$ s
Time Step	2000 s

Two simulations were performed:

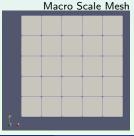
- High Fidelity Simulation
- Macro Scale Simulation using Karst Index

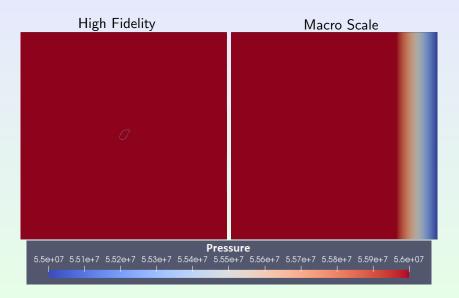
Simulation	High Fidelity	Macro Scale
Number of Elements	578412 (triang.)	25 (quad.)
Number of Nodes	290107	36
Processing time	7 minutes	1.42 seconds

High Fidelity Mesh (zoom)

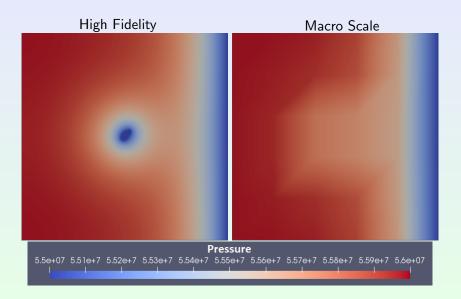
(LNCC)

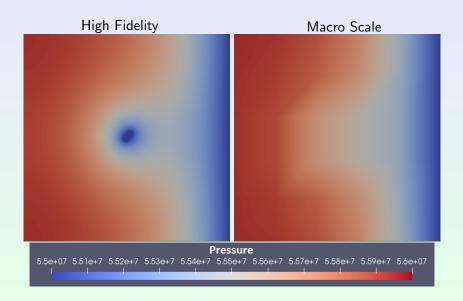




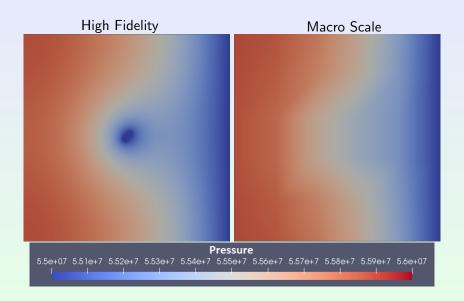




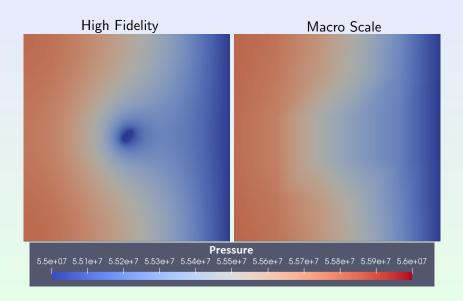


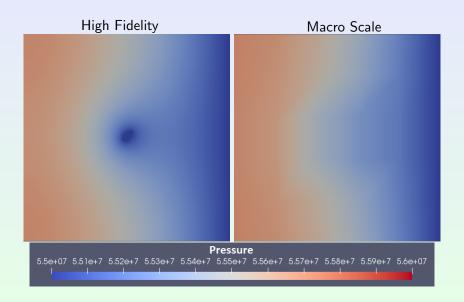




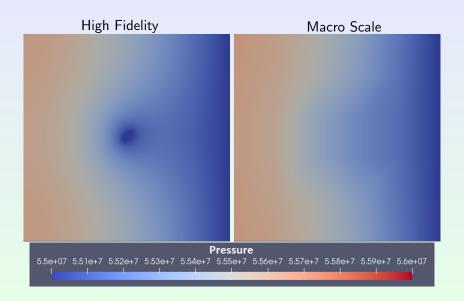




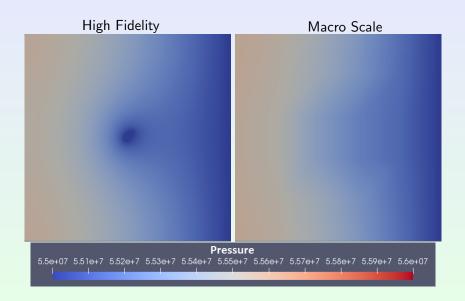




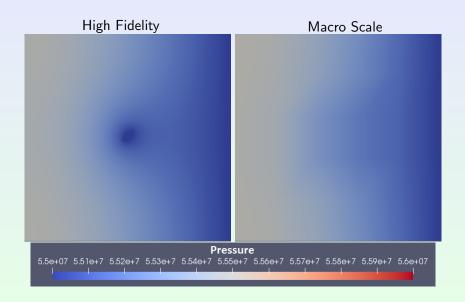


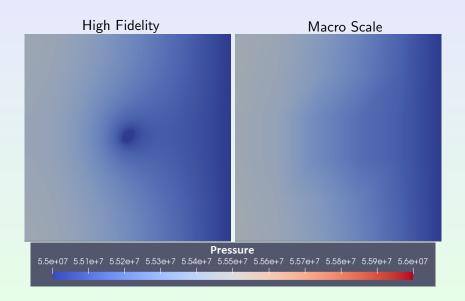


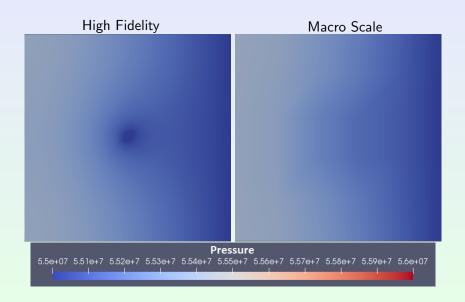


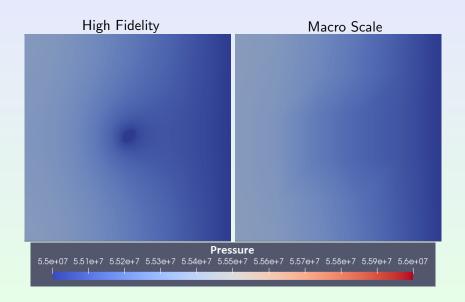


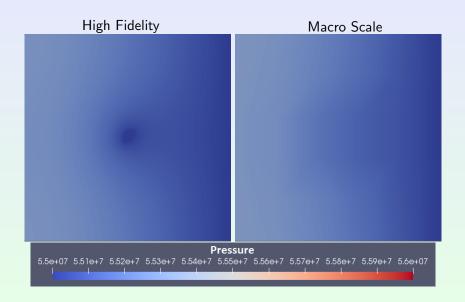
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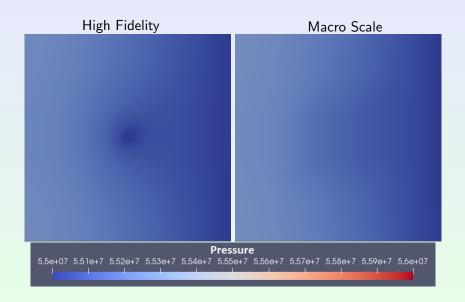


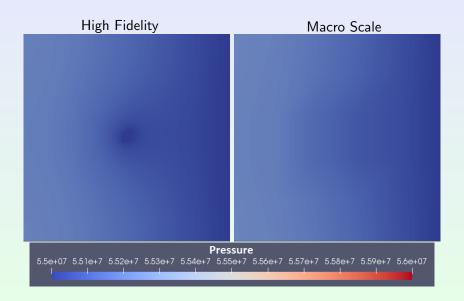


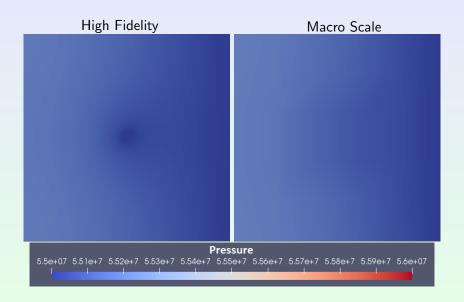


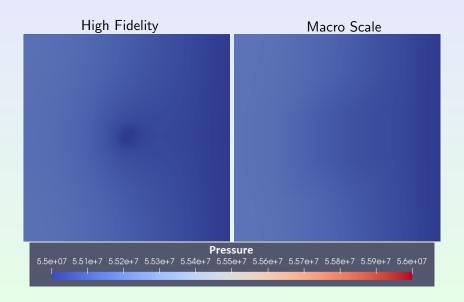


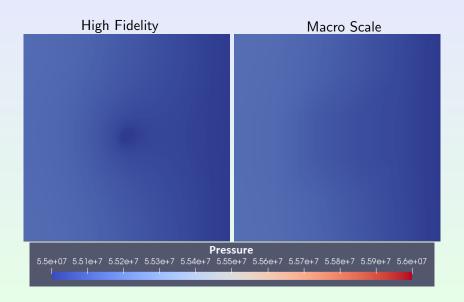




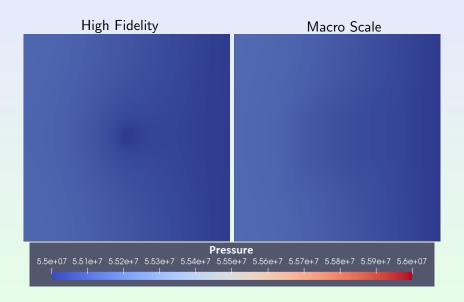




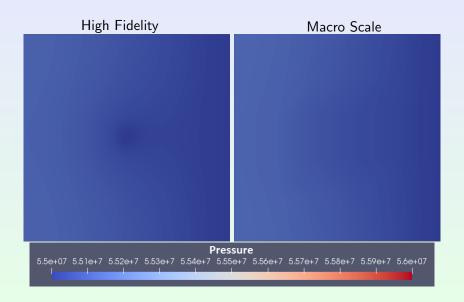




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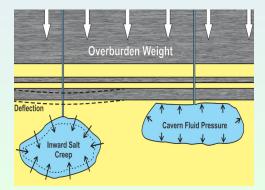


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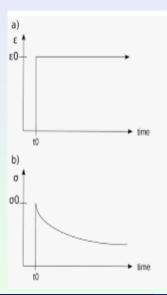
Storage in Salt Dome Caverns

Efficient underground repository

- Inject Fresh Water; withdraw brine dissolving the salt
- Create a stable large Cavity. Salt is chemically inert to CO₂
- Also potental sites for energy storage using highly compressible gases
- Creep. Salt tends to fill the holes gradually moving towards the cavern

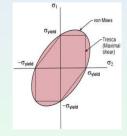


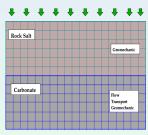
Creep Model for Halite



Geomechanics: Von Mises Stress - Distortional Energy

$$\sigma = -\rho \mathbf{I} + \mathbf{s}, \qquad Q = \sqrt{\frac{3}{2} s_{ij} s_{ji}} \qquad \frac{\partial \mathcal{E}_c}{\partial t} = \epsilon_0 \left(\frac{Q}{\sigma_0}\right)^N \sqrt{\frac{3}{2}} \frac{\mathbf{s}}{\|\mathbf{s}\|}$$





Viscoelastic Model for the Rock Salt

Q – Energy that triggers creep in the saline cap rock

ABC

COMOHR/LNCC

25 / 25

CONCLUDING REMARKS

- Great Opportunity for CO2-storage in Pre-Salt
- Salt Layer
 - Self-Healing due to Creep
 - Chemically Inert to CO2
 - Impermeable Geological Formation

Energy: WAG - Water Alternate Gas Injection for OII Recovery

